



The challenges and need for multi-model intercomparisons of chemistry climate models in the troposphere

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CCM Intercomparisons



- Chemistry Climate Model initiative (CCMi) grew out of the Chemistry Climate Model Validation (CCMVal) activity
 - stretches back to early 2000s when CCMs first began performing long transient simulations
 - these activities have always had as a focus the provision of CCM projections for the WMO/UNEP Ozone Assessment
 - considerable focus on model assessment and 'process-based' diagnostics





- some early-generation CCMs showed unphysically large mixing ratios of Cly
 - very evident problems of mass conservation now much reduced



- CCM chemistry-climate model initiative
- many models show a too late breakdown of the Antarctic vortex
 - delayed transition to easterlies
 - maintains Antarctic ozone hole too late into spring



Average date of zonal average wind at 60°S passing through zero, transitioning from westerly to easterly, in the CCMVal-2 models. Figure 4.2 from SPARC CCMVal (2010).



- McLandress et al. (2012) investigated drag in a version of a middle atmosphere model with data assimilation
 - model drag dominated by orographic
 - gap in drag around 60°S tied to gap in land between South America and Antarctic peninsula





- McLandress et al. (2012) investigated drag in a version of a middle atmosphere model with data assimilation
 - analysis increment for zonal wind add significant extra drag into the area around 60°S
 - free running model with addition of extra orographic GWD around 60°S compared better with observations for temperature, winds vortex breakdown



(a) Delta U (JJA 2006-10)

Zonal average wind tendency (m s⁻¹ day⁻¹) due to analysis increments. Figure 1 from McLandress et al., (2012).



- source of extra drag around 60°S is still unclear
 - orographic waves from small islands (e.g. Hoffmann et al., 2016)
 - secondary gravity waves generated by breaking orographic waves (e.g. Satomura & Sato, 1999)
 - waves generated by winter storms over southern oceans (e.g. Hendricks et al., 2014)
 - horizontal advection and meridional refraction of GWs (e.g. Sato et al., 2012)
 - GWD schemes in models typically assume GWs move vertically
- part of the motivation for the Southern Hemisphere Transport, Dynamics, and Chemistry–Gravity Waves (SOUTHTRAC-GW) mission in 2019
 - aircraft observations of tip of South America with lidar and limb imager



- SOUTHTRAC-GW aircraft observations and modelling show evidence for refraction
 - importance of other processes has not been ruled out



Temperature anomalies at 27 km on September 12 2019, 06Z, from a 3 km WRF simulation and reconstructed from the airborne lidar observations. Figure 6 from Geldenhuys et al. (2023).



- methane lifetime to loss by reaction with OH in the troposphere is nicely constrained by observations of other species
 - methyl chloroform was controlled under the Montreal Protocol with emissions dropping to near zero in the late 1990s
 - accounting for the relative rates of reaction with OH, methane lifetime estimated at 11.2 +/- 1.3 years for reaction with OH (Prather et al., 2012)



Monthly average near-surface concentrations of methyl chloroform in different latitude bands derived from NOAA Global Monitoring Laboratory observations. See Montzka et al. (2011).



• CCMs show a wide range of lifetimes (obs. 11.2 +/- 1.3 years)



The methane lifetime to reaction with OH for the Atmospheric Chemistry and Climate Model Intercomparison (ACCMIP) models. Figure 1 from Voulgarakis et al. (2013).



The annual cycle in methane lifetime to reaction with OH for the CCMi-1 models and additional chemical transport models for year 2000 conditions. Figure 2 from Nicely et al. (2020).

- Nicely et al. (2020) created neural networks to reproduce the 3-D distribution of OH in e a selection of input va
 - swapping inputs from to another one to inve sensitivity of OH to dif

Original CH₄ lifetim **OsloCTM** for Janua

Change in CH₄ lifetime using th as input to the GMI emulator a

Table 1. Accounting of CH₄ lifetime differences between GMI and OsloCTM simulations for January 2000.

OH in each model from			GMI	OsloCTM
nput variables	$\tau^{a}_{CH_{4},ORIG}$ (year)		9.24	7.18
outs from the native model ne to investigate the OH to different factors	$\Delta \tau_{CH_4}$ due to ^b :	O_3	-0.91	+0.79
		HCHO	-0.59 -0.64	+0.60 +0.51
		NO_x JNO ₂	$-0.45 \\ -0.34$	+0.33 +0.15
CH ₄ lifetime in GMI and for January 2000		Isoprene	-0.03 +0.19	+0.28 -0.07
		H_2O	+0.19 $+0.10$	-0.13
e using the OsloCTM HCHO mulator and vice versa.		CH_4^{NOKM} NO/NO _x	+0.11 +0.07	-0.06 -0.05
		$O_3 \text{ COL}$ T	$-0.02 \\ -0.02$	-0.06 + 0.00
Residual	$\Delta \tau^{c}_{CH_4,TOT}$		-2.52	+2.30
	$\tau_{CH_4,ORIG} + \Delta \tau_{CH_4,TOT}$ Mech. ^d		$\frac{6.71}{+0.47}$	9.48 -0.24
Table 1 from Nicely et al. (2020).				





Residual

 summary of changes across all possible permutations of input fields from different models into the emulator of each model



Figure 5 from Nicely et al. (2020).



- CCMi-1 specified two types of historical hindcast
 - REF-C1 free running model with specified historical SSTs / sea-ice
 - REF-C1SD historical SSTs / sea-ice plus nudging of dynamics to reanalysis
- a number of synthetic tracers in the troposphere to estimate transport timescales in the models
 - specified tracer concentration (100 ppb) at the surface for all points between 30°N to 50°N
 - concentration decays with a specified lifetime of 5 days (T_5) or 50 days (T_{50})



- Orbe et al. (2018) showed that the climatological distribution in mid-troposphere shows a large range across models
 - tied to a large range in the parameterized convective updraft mass flux



The 2000 – 2009 seasonal average zonal average concentrations over 700 – 400 hPa for the 5- and 50day idealized tracers. Figure 2 from Orbe et al. (2018).

Conclusions



- by highlighting aspects with considerable model diversity or common biases, MIPs have the ability to highlight model shortcomings or uncertainties
- longer history of model intercomparison / assessment in the stratosphere
 - some progress on understanding processes and improving models
 - the role of mixing in age of air is better understood, even though issues remain
- Nicely et al. ML analysis of CCMi-1 models showed a number of factors are important in explaining the model variability for OH / CH4 lifetime
 - J[O3-> O(1^D)], O₃, NO_x, H₂O, CO, NO/NO_x ratio, HCHO, residual (unidentified) factors
- Orbe et al. analysis showed a range (2X 3X) of transport timescales from boundary layer to free troposphere and to the Arctic
- likely a lot of linkages between these two findings

Conclusions



- the state of the atmosphere is underconstrained by observations and the troposphere is such a 'messy' place, we need new and thoughtful ways to analyse processes in models
 - can we still make progress asking for monthly average fields from models
- for a process like vertical transport by deep convection, what is 'correct'
 - is there a role for convection resolving global models as a benchmark?
- we are (just) beginning to think of a new phase of CCMI with a focus on the troposphere and we welcome participation by people with ideas for new types of analyses, diagnostic fields, process studies, observations to compare against,