Relationship between Water Vapor and Cold Point Tropopause during Boreal Summer

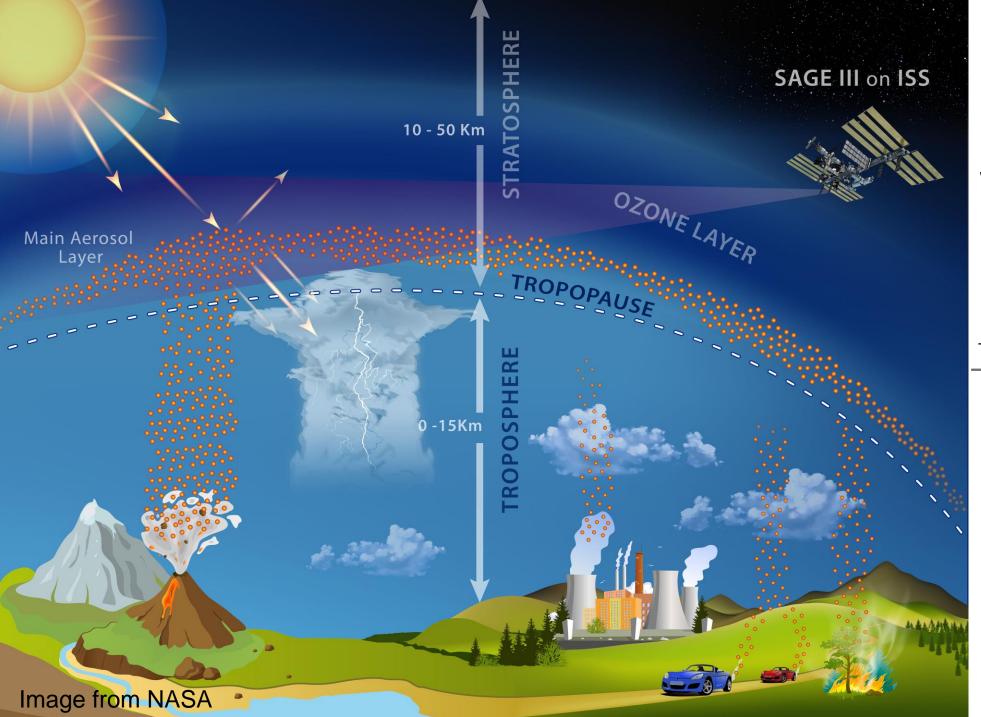


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Mijeong Park ACOM, NSF NCAR, Boulder, USA June 3-7, 2024 STIPMEX, Pune, India

What I am going to talk about today.

- 1. Water vapor over Asian vs. N. American monsoon regions
- 2. Cold point tropopause (CPT) in the UTLS region
- 3. COSMIC-2 temperatures and atmospheric waves
- 4. Plans In situ measurements and a global chemistry climate model (WACCM)



STRATOSPHERE

TROPOPAUSE (8-16 KM)

TROPOSPHERE

Stratospheric H₂O is an important driver of decadal global surface climate change.

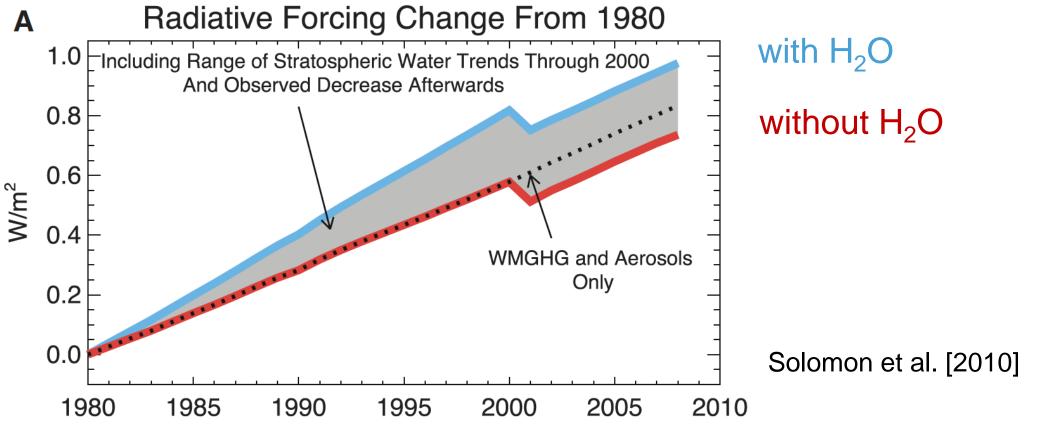
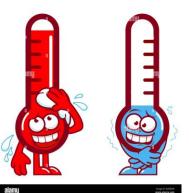


Fig. 3. Impact of changes in stratospheric water vapor on surface climate. (A) Time series of the changes in radiative forcing since 1980 due to well-mixed greenhouse gases (WMGHG), aerosols, and stratospheric water vapor.

Increase in stratospheric H_2O will accelerate the warming of surface temperature.



Lower stratospheric water vapor impacts atmospheric circulation.

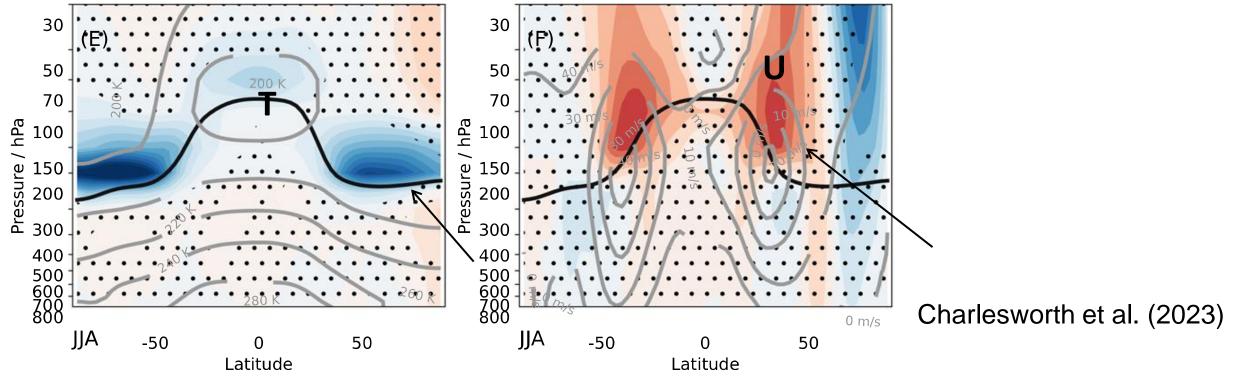
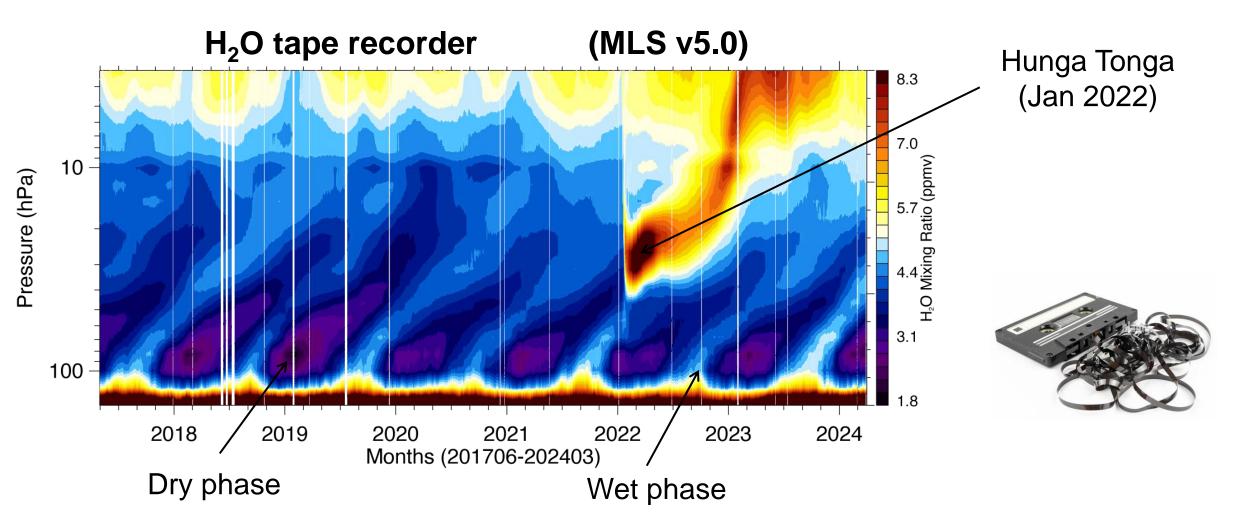


Fig. 2 Atmospheric circulation effects and dynamical mechanism induced by lowermost stratospheric water vapor changes. (E) Temperature, (F) Zonal wind (JJA).

Upward and poleward shift of subtropical jets Strengthening of the stratospheric circulation The seasonal variation in H_2O entering the stratosphere follows the CPT temperature, setting the "base" for the tropical tape recorder during the dry phase.



During the wet phase of the tropical tape recorder, enhanced H₂O over the NH summer monsoons makes its way into the lower stratosphere.

H₂O in the Stratosphere – Boreal Summer (JJA)

Randel et al. (2012)

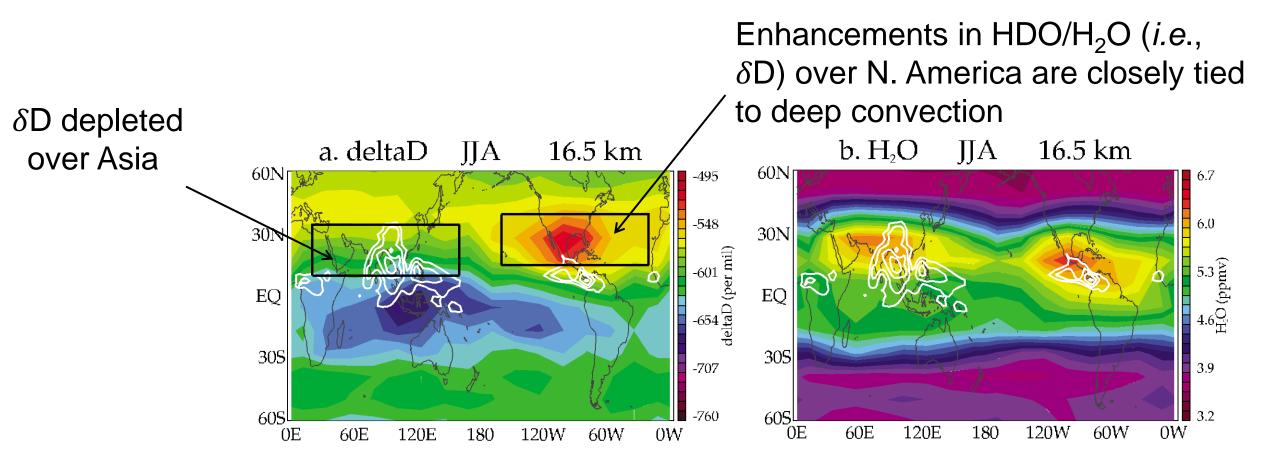
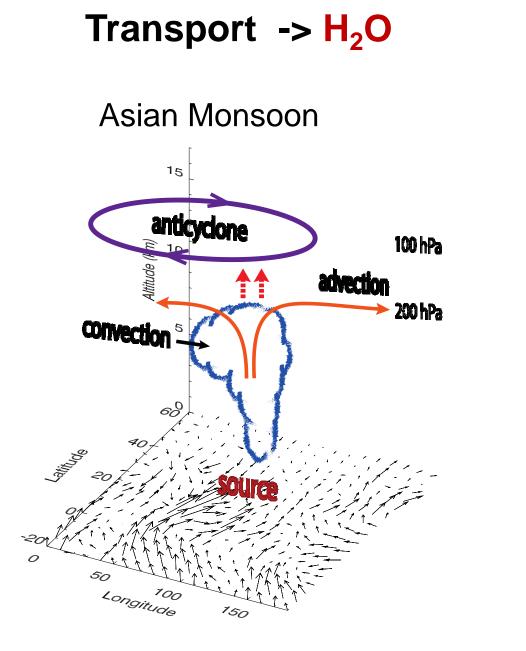


Figure 12. Cross sections of (a) dD and (b) H₂O at 16.5 km during JJA. The dD field has been corrected for methane effects, although this effect is small and only influences the highest latitudes. White contours denote strongest climatological tropical convection, and the black boxes denote the averaging regions for Figures 13 and 14. Note the isotopic enrichment correlated with high water vapor over the North American (NA) monsoon but the lack of a similar signal over the SE Asian monsoon region.

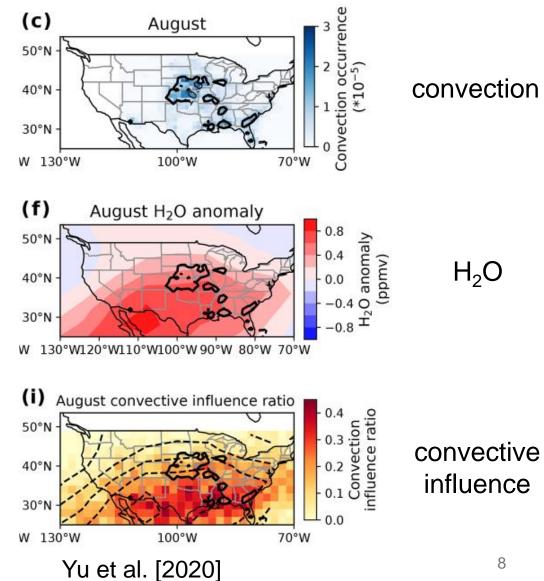
 $\delta D(\text{\%}) = \left\{ \left[(HDO/H_2O)/(HDO/H_2O)_{SMOW} \right] - 1 \right\} \times 1000,$



Park et al. [2009] Ploeger et al. [2013]

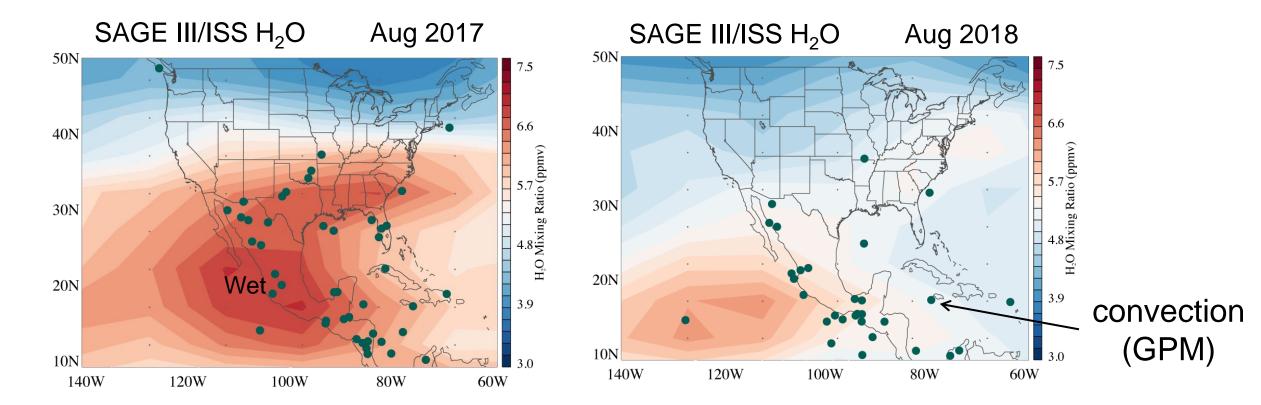
Convection -> H_2O

North American Monsoon



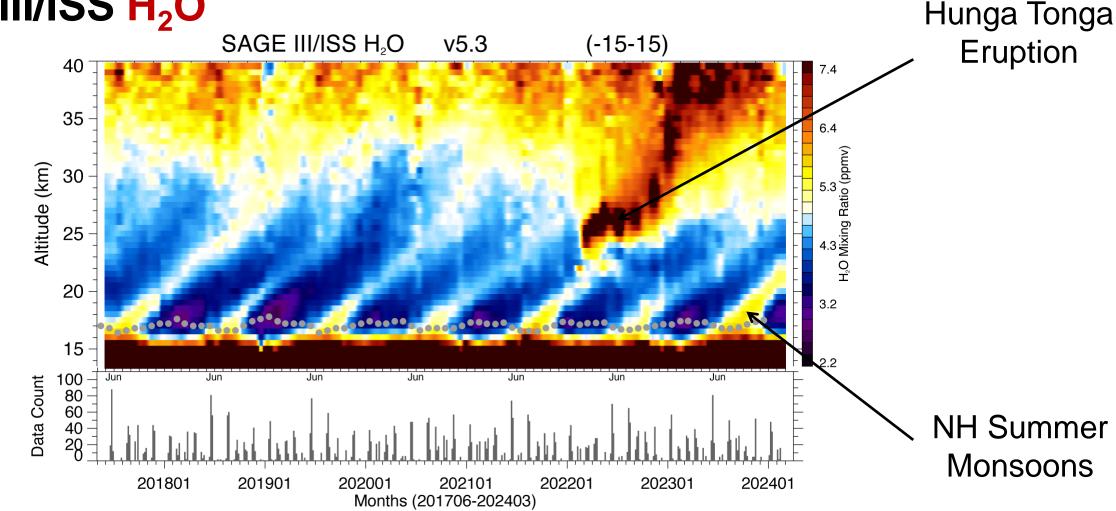
Convective influence in H₂O?

N. American Monsoon



Convective influence on lower stratospheric water vapor is unclear.

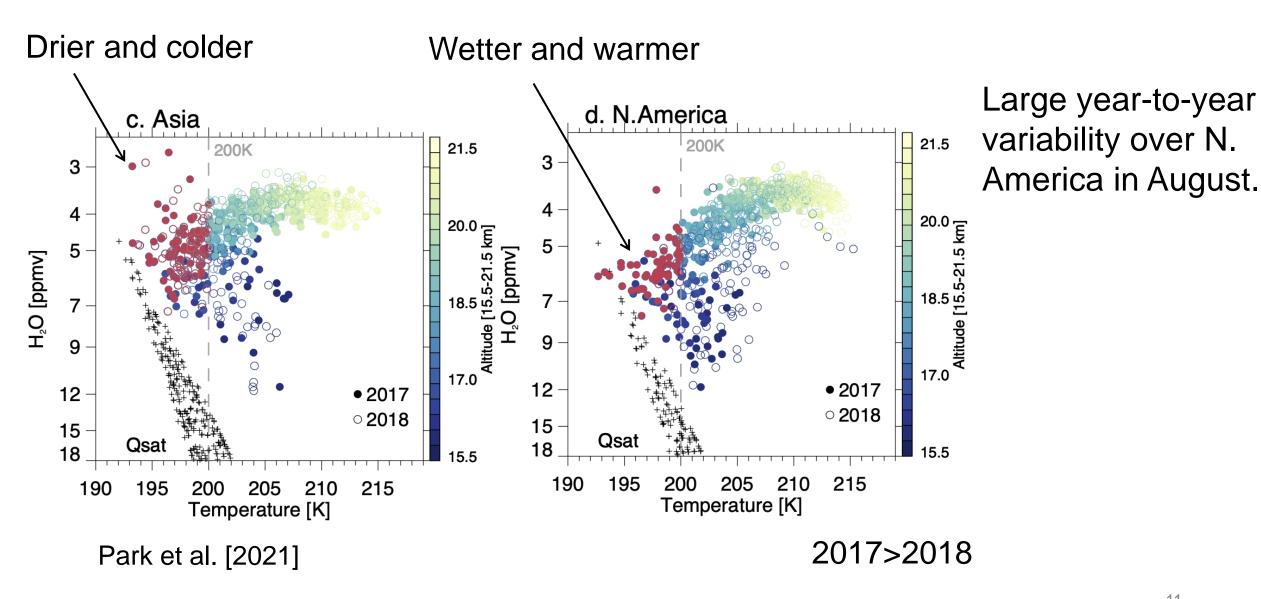
SAGE III/ISS H₂O



[SAGE III/ISS]

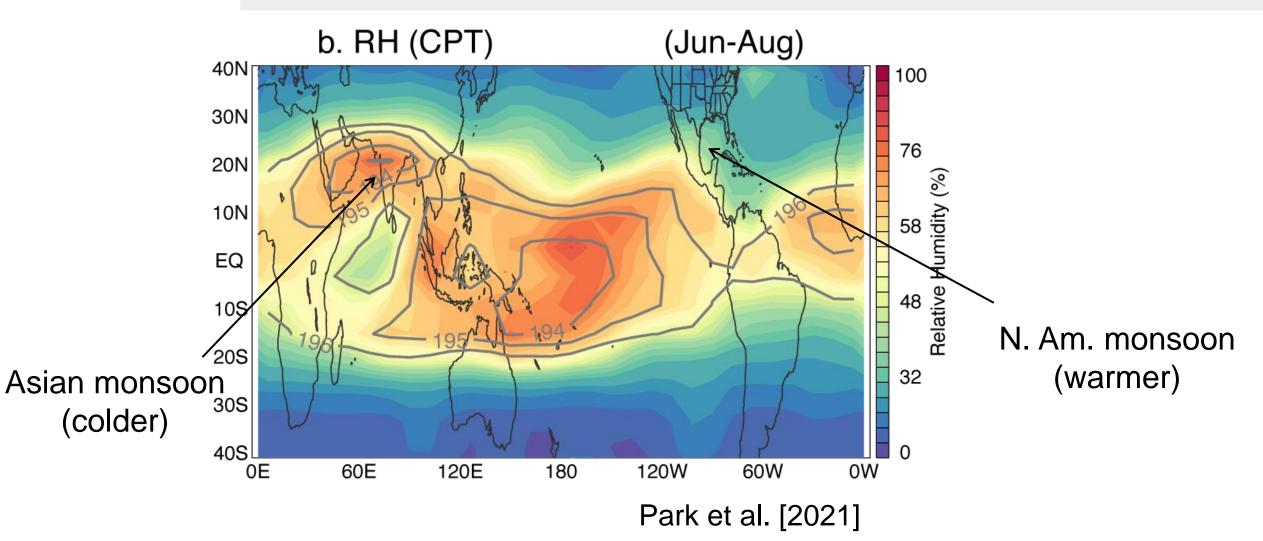
The Stratospheric Aerosol and Gas Experiment III on the International Space Station (SAGE III/ISS) was launched in **February 2017**. The SAGE III/ISS instrument provides measurements of aerosol extinction, NO₂, O₃, and H₂O between 70°S-70°N latitude using the techniques of solar occultation.

Asian vs. N. American monsoon regions (2017 vs. 2018)



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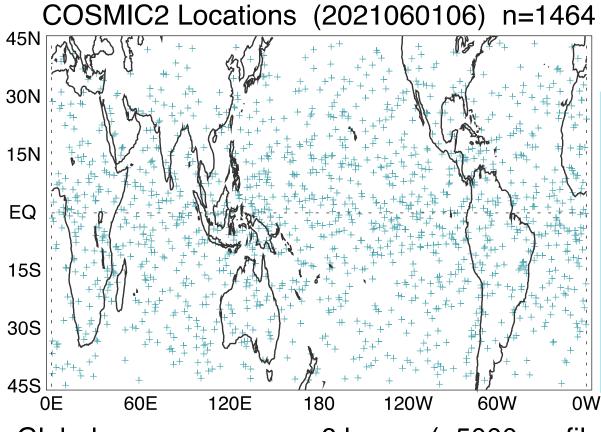
 $H_2O - CPT$ The enhanced H_2O over the Asian monsoon region is collocated with colder CPT temperature. In comparison, tropopause over the North American monsoon is warmer.



COSMIC-2 Temperature

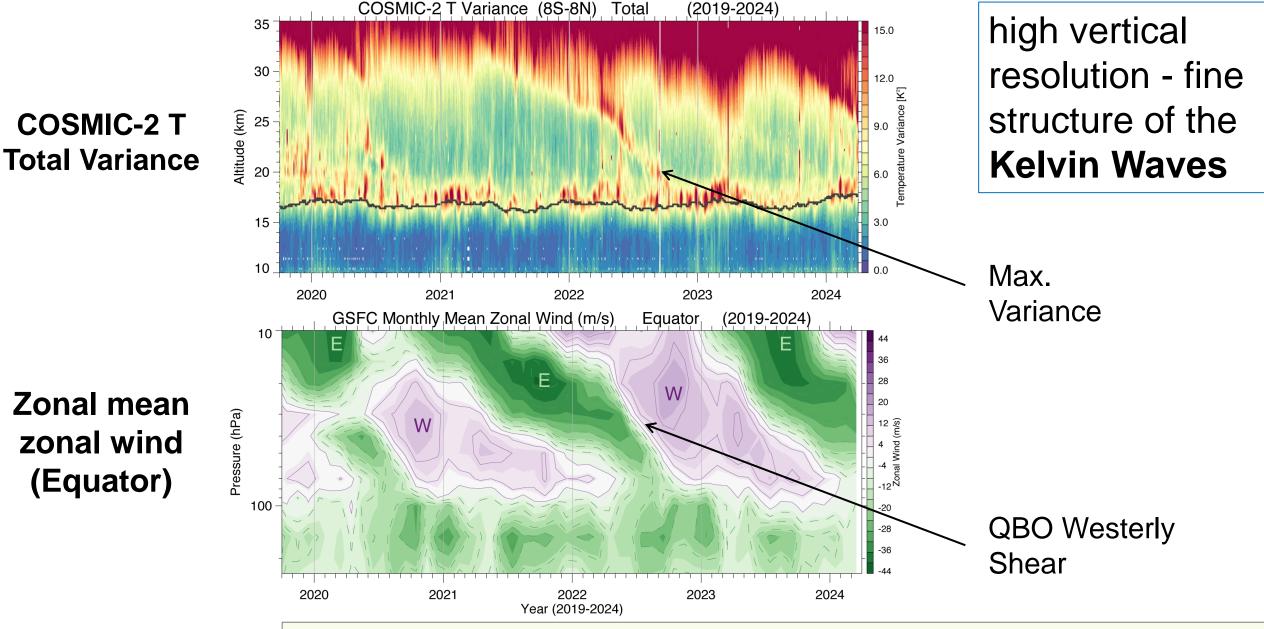
The Constellation Observing System for Meteorology, Ionosphere and Climate-2 is an equatorial constellation of radio occultation satellites that was launched on **25 June 2019.**





COSMIC-2 provides key observations for characterizing **equatorial waves**. An important aspect of equatorial waves is their relatively narrow vertical scales (typical vertical wavelengths of \sim 4–8 km).

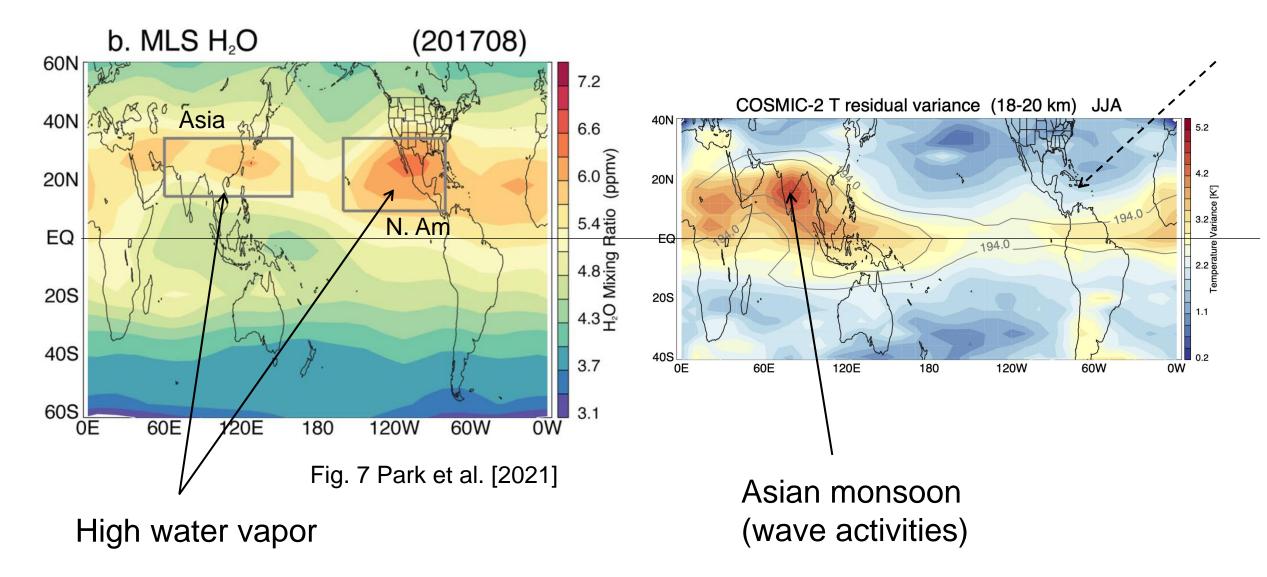
Global coverage - every 6 hours (~5000 profiles/day)



The quasi-biennial oscillation (**QBO**) dominates the variability of the equatorial stratosphere (~16–50 km) and is seen as downward propagating easterly and westerly wind regimes, with a variable period averaging ~ 28 months [Baldwin et al., 2001].

H₂O vs. Temp Variance (waves)

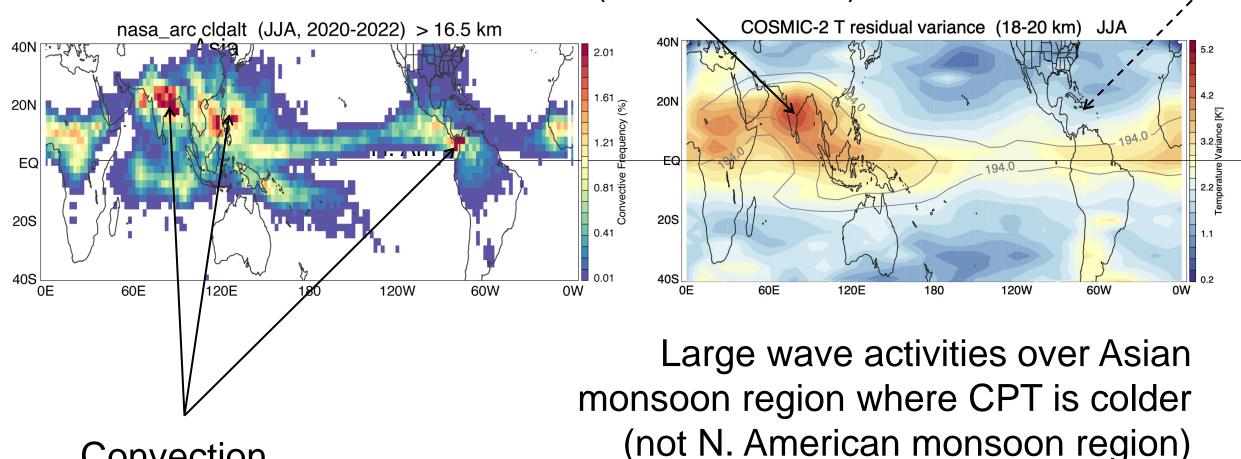
Asia vs. N. America



Convection *vs.* **Temp Variance** (waves)

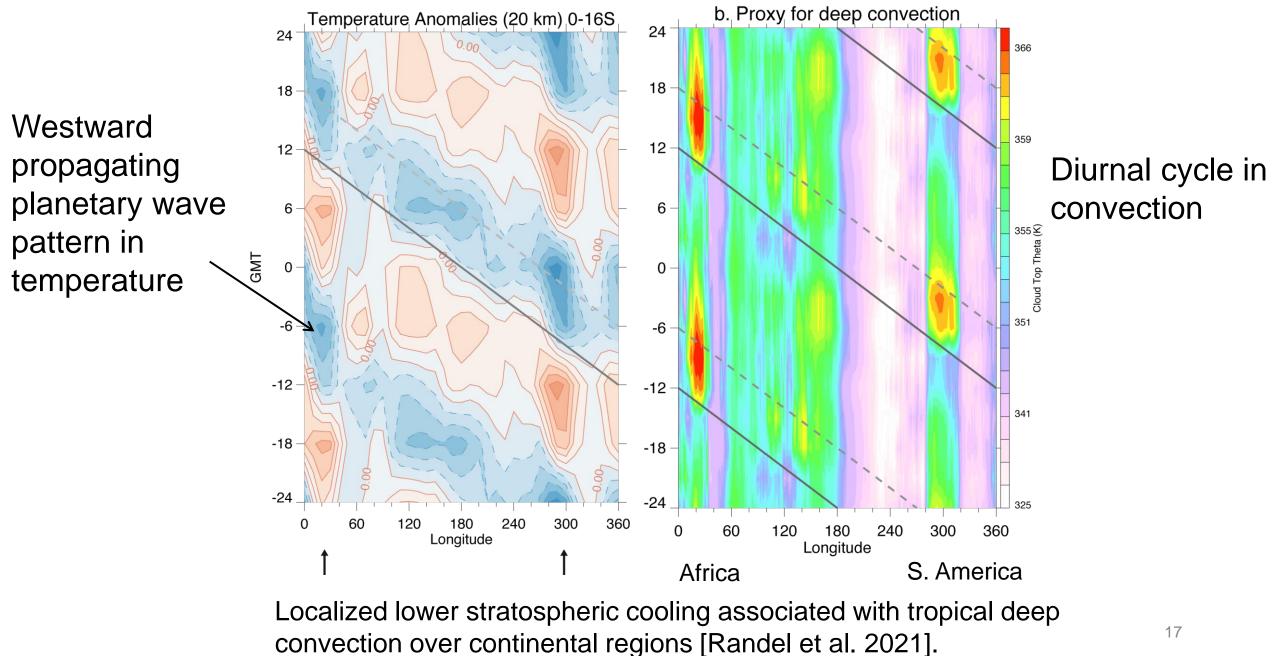
Asia vs. N. America

Asian monsoon (wave activities)

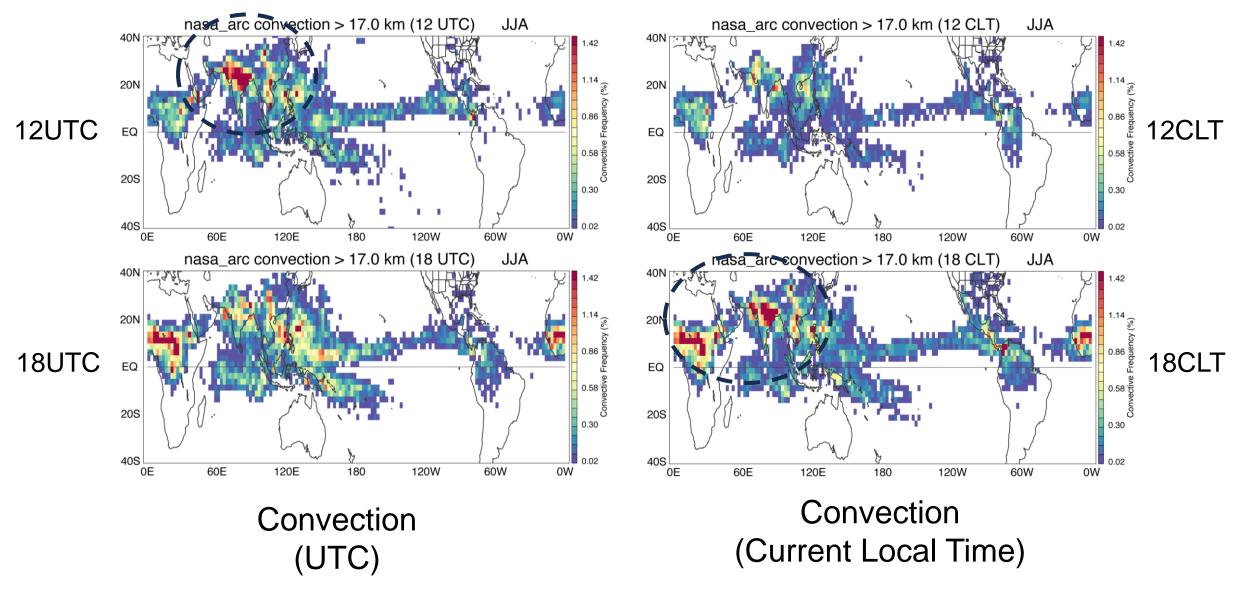


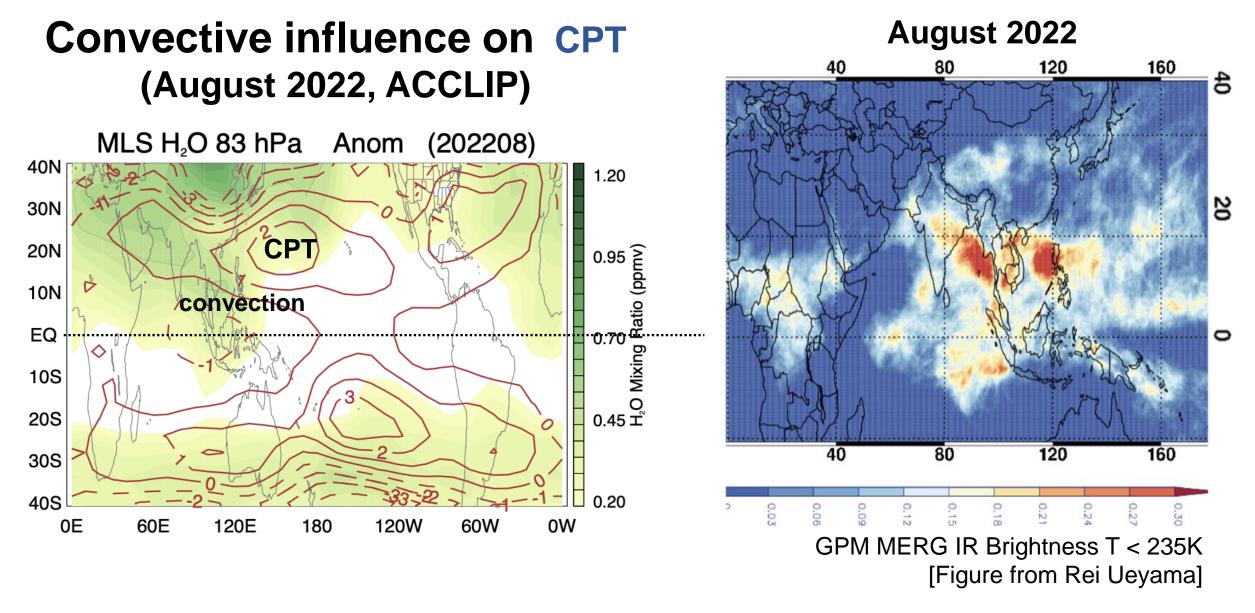
Convection

Deep convection modifies the temperature (diurnal cycle).



Diurnal cycle in convection better represented in current local time.





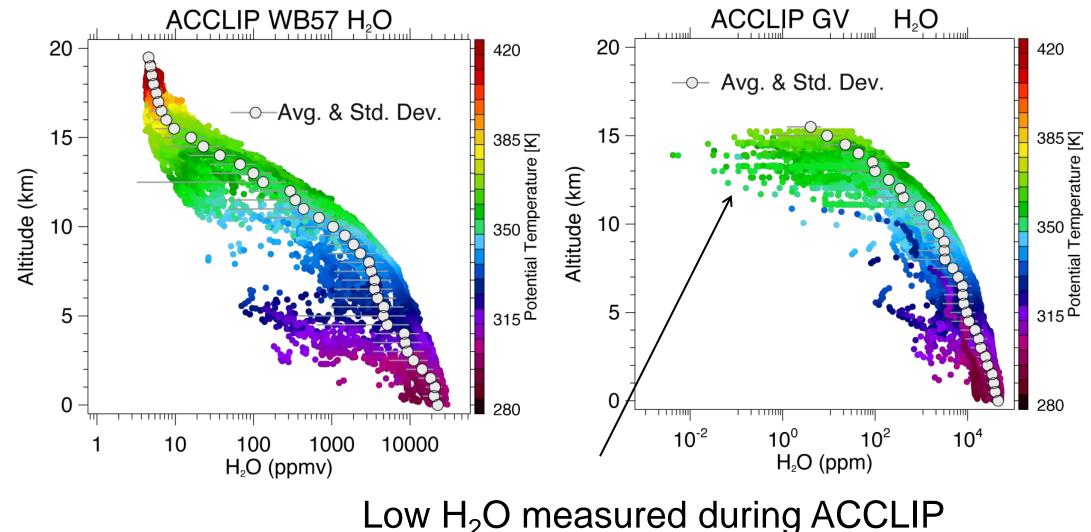
Convection was more frequent over Bay of Bengal and west of the Philippines in August 2022 compared to climatology (2006-2022).

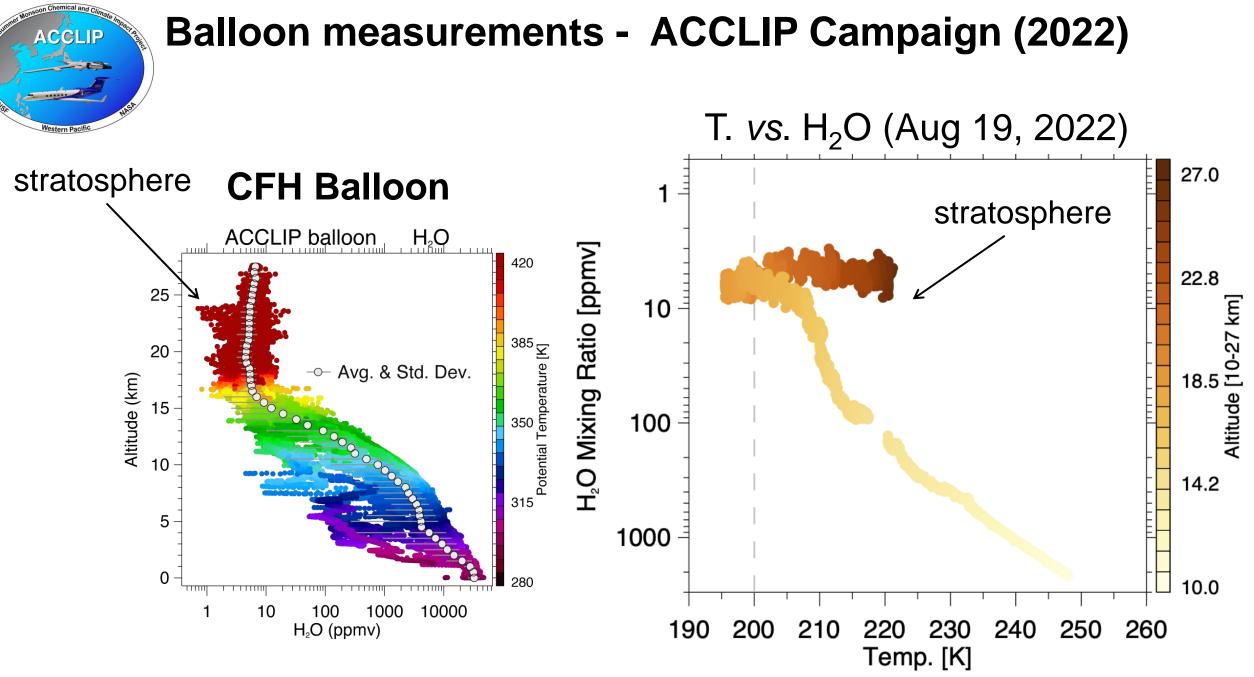
In situ measurements - ACCLIP Campaign (2022)

NASA WB57

ACCLIP

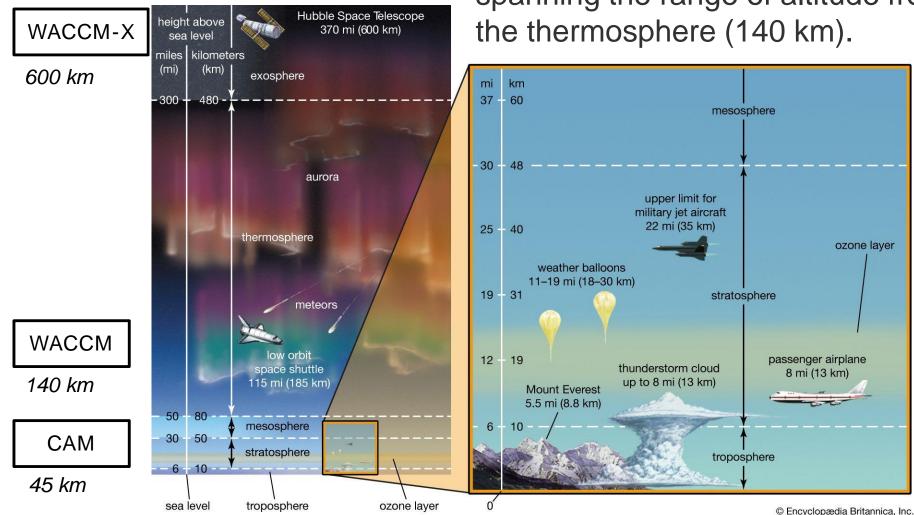
NSF GV





Further investigate Temperature & H₂O relationships over Asia.

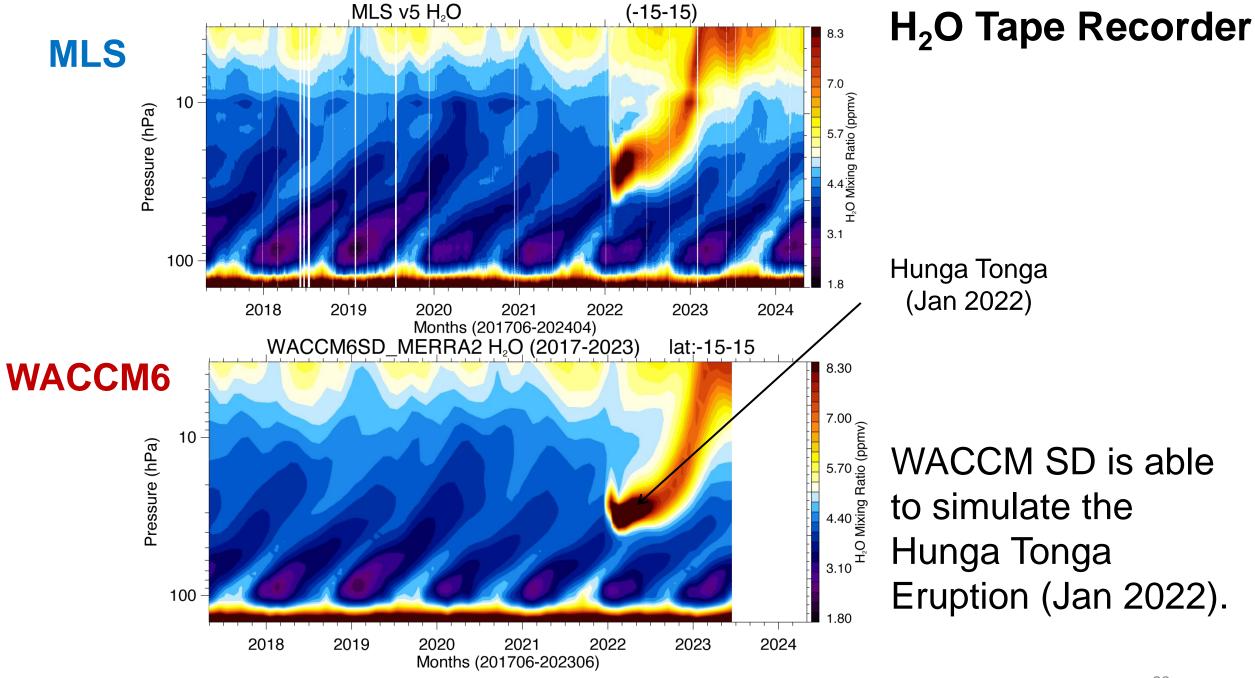
Global chemistry climate model (WACCM) simulations



The Whole Atmosphere Community Climate Model (**WACCM**) is a comprehensive numerical model, spanning the range of altitude from the Earth's surface to the thermosphere (140 km).

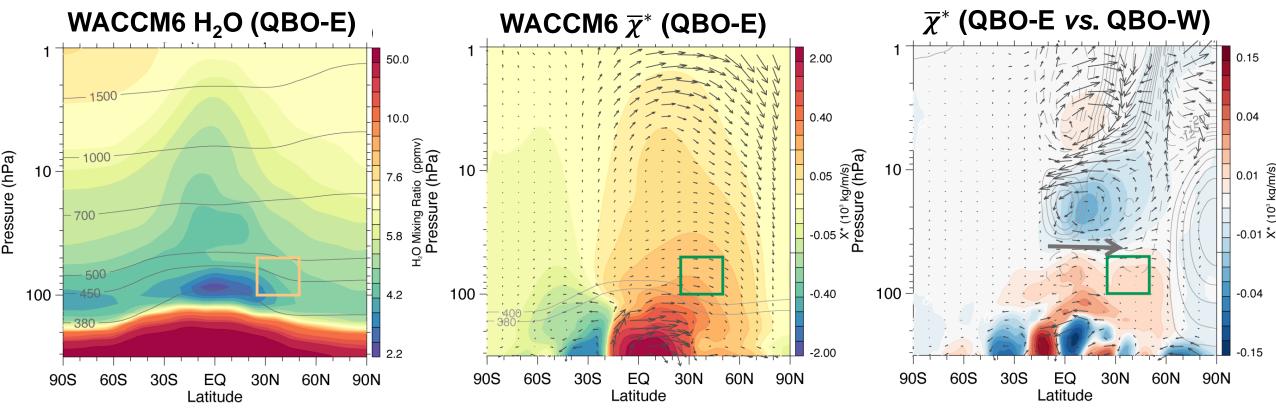


https://cdn.britannica.com/42/90442-050-6CB42E65/layers-atmosphere-Earth-phenomena-heights.jpg



WACCM SD simulations (Jun Zhang)

Changes in H₂O and circulation in WACCM (e.g., QBO)



Transformed Eulerian Mean (**TEM**) Mass stream function $(\bar{\chi}^*)$ Positive: clockwise Negative: counter-clockwise Stronger transport from the tropics to midlatitude during QBO-E.

Summary & Plans

- 1. During the NH summer, stratospheric H_2O is enhanced over Asian and North American monsoon regions in the UTLS. However, the correlations between H_2O , CPT and convection are unclear.
- Large-scale transport plays a role in controlling lower stratospheric H₂O in boreal summer. Convection might control water vapor indirectly by modifying atmospheric waves thorough affecting local temperatures.
- Higher H₂O and relatively warmer CPT were observed over the Western Pacific during the ACCLIP campaign in August 2022. Further work is needed to examine the correlations between H₂O and CPT.
- Along with observational datasets, a global chemistry climate model (WACCM) will provide quantitative information on the changes in largescale circulation in the UTLS region.